

## Heating of the atmosphere by the sea

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**Abstract.** An important component of atmospheric heating is the the surface heat flux from the sea. We give expressions for the main components of this flux.

The heat flux components are latent, sensible and radiative heating. These can be parameterised in terms of the surface temperature  $T_s$  as follows.

### 1. Latent Heat

$$Q_L = L\rho K_L |V| (q_s - q_a)$$

where  $L$  is the latent heat of vaporisation of water,  $\rho$  is the air density,  $K_L$  is an empirical constant of order unity,  $V$  is the wind velocity,  $q_a$  is atmospheric specific humidity,  $q_s$  is the specific humidity corresponding to sea-surface temperature  $T$  and is given by

$$q_s = \frac{0.622e}{p - 0.378e},$$

where  $e$  is the vapour pressure in millibars given by

$$\log_{10} e = 9.4 - \frac{2.35 \times 10^3}{T_s}.$$

### 2. Sensible heat

$$Q_s = C_p \rho K_s |V| (T_s - T_a)$$

where  $C_p$  is the specific heat of air at constant pressure,  $\rho$  is the density of air,  $K_s$  is an empirical constant of order unity, and  $T_a$  is air temperature.

### 3. Net longwave radiation

$$Q_p = Q_0(1 - 0.62C + 0.0019\alpha)(1 - A) - \varepsilon\sigma T_a^4(0.39 - 0.05e^{0.5})(1 - a_0C^2) + 4\varepsilon\sigma T_a^3(T_s - T_a)$$

where  $Q_0$  is the clear sky radiation received at the surface,  $C$  is the cloud cover fraction in tenths,  $\alpha$  is the solar noon altitude,  $A$  is the oceanic albedo,  $\varepsilon$  is the emissivity of water,  $\sigma$  is the Stefan-Boltzmann constant and  $a_0$  is a constant.